



# On the formation of diaplectic glass: Shock and thermal experiments with plagioclase of different chemical compositions

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(Received 24 April 2018; revision accepted 14 March 2019)

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**Abstract**—This contribution addresses the role of chemical composition, pressure, temperature, and time during the shock transformation of plagioclase into diaplectic glass—i.e., maskelynite. Plagioclase of An<sub>50-57</sub> and An<sub>94</sub> was recovered as almost fully isotropic maskelynite from room temperature shock experiments at 28 and 24 GPa. The refractive index (RI) decreased to values of a quenched mineral glass for An<sub>50-57</sub> plagioclase shocked to 45 GPa and shows a maximum in An<sub>94</sub> plagioclase shocked to 41.5 GPa. The An<sub>94</sub> plagioclase experiments can serve as shock thermobarometer for lunar highland rocks and howardite, eucrite, and diogenite meteorites. Shock experiments at 28, 32, 36, and 45 GPa and initial temperatures of 77 and 293 K on plagioclase (An<sub>50-57</sub>) produced materials with identical optical and Raman spectroscopic properties. In the low temperature (<540 K) region, the formation of maskelynite is entirely controlled by shock pressure. The RI of maskelynite decreased in heating experiments of 5 min at temperatures of >770 K, thus, providing a conservative upper limit for the postshock temperature history of the rock. Although shock recovery experiments and static pressure experiments differ by nine orders of magnitude in typical time scale (microseconds versus hours), the amorphization of plagioclase occurs at similar pressure and temperature conditions with both methods. The experimental shock calibration of plagioclase can, together with other minerals, be used as shock thermobarometer for naturally shocked rocks.

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## INTRODUCTION

The term diaplectic glass (from the Greek word diaplesso = to destroy by striking) was introduced by von Engelhardt et al. (1967) to describe an amorphous phase without vesicles and flow structures. It retains the grain boundaries, chemical composition, and zonation of the primary mineral. The density and refractive indices (RI) (a dimensionless number describing the speed of light in a material) of diaplectic glasses are lower than those of the primary crystal and

higher than those of a quenched mineral glass (Bunch et al. 1967; von Engelhardt et al. 1967; Osinski et al. 2018). An amorphous phase forming pseudomorphs after plagioclase (NaAlSi<sub>3</sub>O<sub>8</sub>–CaAl<sub>2</sub>Si<sub>2</sub>O<sub>8</sub>) was first described in the Shergotty achondrite by Tschermak (1872, 1883), and the author named this phase maskelynite, after the British mineralogist Mervyn Herbert Nevil Story Maskelyne. Milton and De Carli (1963) showed that maskelynite can be produced from plagioclase with bytownite composition (An<sub>80</sub>) by shock recovery experiments at pressures of 25–35 GPa and

calculated postshock temperatures of 490–590 K. Binns (1967) considered the lack of eutectic or eutectic-like melting phenomena with adjacent mafic minerals as an argument against a purely thermal origin of maskelynite. Instead, such diaplectic glasses form as a result of the material interacting with shock waves that are sufficiently intense to destroy the crystal lattice. After unloading the pressure, the material is left at a temperature below its melting point (von Engelhardt et al. 1967). Diaplectic glass is an umbrella term for shock amorphization including the open framework tectosilicates quartz and feldspar.

There is still a substantial discussion about the mechanism as well as the pressures and temperatures needed to form diaplectic glass during natural impact events, with pressures ranging between  $<22$  and  $>80$  GPa and temperatures varying by more than 1000 K (Stöffler et al. 1986; Sharp et al. 1999; Chen and El Goresy 2000; Kubo et al. 2010; El Goresy et al. 2013). A “hot melt” formation of diaplectic glasses assumes shock melting and subsequent quenching to a dense glass while still at elevated pressures. This process was proposed to explain the development of planar deformation features (PDF) in quartz (Grady 1980). Langenhorst (1994) and Stöffler and Langenhorst (1994) argued that above 35 GPa the total quartz crystal is engulfed in molten PDF and that diaplectic glass represents a quenched high pressure melt. This model was later used to explain the formation of maskelynite (Chen and El Goresy 2000; El Goresy et al. 2013). A “cold” solid state amorphization and mechanical deformation of plagioclase was considered for the formation of maskelynite based on petrographic observation in meteorites and shock and heating experiments (e.g., Duke 1968; Hörz and Quaide 1973; Stöffler et al. 1986; Fritz et al. 2005a; Jaret et al. 2015). A solid state amorphization was observed after static pressure experiments at controlled temperatures of 293 K on Ca-rich ( $An_{96}$ ) plagioclase (Williams and Jeanloz 1989) and on quartz (Hemley et al. 1988; Kingma et al. 1993) and during pre-cooled shock experiments at initial temperatures of 77 K on quartz (Fritz et al. 2011). Such a solid state amorphization is essentially consistent with the gradual formation of diaplectic glasses via mechanical disaggregation of the crystal lattice into domain sizes so small that they are beyond coherent X-ray diffraction (Hörz and Quaide 1973). In the present study, shock experiments on plagioclase of different chemical composition, and subjected to different pressures and initial temperatures, including annealing experiments on the resulting maskelynites are reported. These results are used together with shock and static pressure experiments literature data to discuss the influence of shock

pressure, temperature, time, and composition on the formation of diaplectic glasses, in the present case maskelynite. This aids in resolving controversies on how to interpret the shock metamorphic record in meteorites (for a recent review, see Fritz et al. 2017).

## METHODS

Samples of natural rocks containing plagioclase of different composition were experimentally shocked. Rock disks of 15 mm diameter and 0.6 mm thickness were cut from a troctolite hand specimen B59-36-A of the Gills Allard Economic Geology Collection, University of Georgia Museum of Natural History, Athens, Georgia, USA, and from a gabbro hand specimen from Grass Valley, California, USA, containing plagioclase with a chemical composition of  $An_{50-55}$  and  $An_{94}$ , respectively. The disks were polished on both sides and placed into a high impedance container made from ARMCO iron. The rock samples and the ARMCO container for the shock experiments were prepared at the Museum für Naturkunde in Berlin, Germany (MfN). The shock recovery experiments using the reverberation technique were conducted at the Ernst Mach Institut für Kurzzeitdynamik (EMI) in Freiburg, Germany following the procedures described in Müller and Hornemann (1969), Meyer et al. (2011), and Fritz et al. (2011). Our shock experiments were conducted in the range of 20–45 GPa and at initial temperatures of 293 K (room temperature) to study the shock behavior of plagioclase with different chemical composition. Additional experiments at 77 K, the boiling temperature of liquid nitrogen, were conducted to document the influence of temperature on the formation of maskelynite. Experimental details are given in Table 1, and the pressure in the sample is considered to be accurate to 4%, i.e.,  $\pm 1.1$ – $1.8$  GPa (Müller and Hornemann 1969).

After conducting the shock experiments, the ARMCO iron containers were machine-opened and the rock disks removed. From the polymineralic rock disks, plagioclase grains were handpicked. The RI of individual grains was determined at the MfN using the immersion method (Fritz et al. 2002, 2005a).

This was followed by heating experiments of individual maskelynite grains that were placed for 5 min into a preheated oven. The temperature of the oven was increased in 50 K temperature heating steps from 570 to 1170 K. Samples were inserted in the oven after it had acquired the new temperature, and the grains were kept at that temperature for 5 min. The RI was determined after each temperature step.

The chemical composition of plagioclase was determined using thin sections of unshocked material and polished grain mounts of single grains for which

Table 1. Conditions of the shock experiments with 0.6 mm thick and 15 mm diameter rock disks.

RS	$P$ (GPa)	$T$ (K)	$C$ (mm)	$D$ (mm)	Expl.	Init. $P$ (GPa)	$V$ (km s <sup>-1</sup> )	$t$ ( $\mu$ s)
Tr (An <sub>50-57</sub> )	28	77	6.2	4	C4 (64)	33	1.62	0.53
Tr	28	293	6.2	4	C4 (64)	33	1.62	0.53
Tr	32	77	8.8	3	C4 (64)	47	2.16	0.1
Tr	32	293	8.8	3	C4 (64)	47	2.16	0.1
Tr	36	77	6.5	3	C4 (64)	47	2.16	0.2
Tr	45	77	6.3	3	Comp. B (64)	52	2.35	0.27
GV (An <sub>94</sub> )	20	293	5.0	4	TNT	30	1.55	0.83
GV	24	293	11.0	4	C4 (64)	33	1.62	0.43
GV	28	293	6.2	4	C4 (64)	33	1.62	0.53
GV	36	293	6.5	3	C4 (64)	47	2.16	0.2
GV	41.5	293	9.5	3	Comp. B (64)	52	2.35	0.16

Type of rock sample (RS) being either Troctolite (Tr) or Grass Valley Anorthite (GV), shock pressure in the sample ( $P$ ), initial temperature of the experiment ( $T$ ), thickness of cover plate ( $C$ ), thickness of driver plate ( $D$ ), type of explosive (Expl.), initial pressure on top of the iron container (init.  $P$ ), velocity of the driver plate ( $V$ ), and pulse length of shock compression ( $t$ ) for shock pressure ( $P$ ).

the RI has been previously determined. Quantitative mineral chemical analyses were obtained at the MfN with a JEOL JXA 8500F field emission cathode electron microprobe equipped with five wavelength-dispersive spectrometers. Minerals were analyzed with a 10–15 kV accelerating voltage, a 15 nA beam current, and a defocused 10  $\mu$ m diameter beam. Suitable glass and mineral standards certified by the Smithsonian Museum were used as reference samples for electron microprobe analysis (Jarosewich et al. 1980).

Raman spectra were obtained at the Deutsches Zentrum für Luft- und Raumfahrt (DLR), Berlin, Germany using a confocal Raman microscope WITec alpha 300 R of the polished grain mounts. The samples were excited with a 532 nm Nd:YAG laser. The Raman scatter was collected using a long working distance 100 $\times$  objective (NA 0.8). The spot size on the sample was smaller than 1.5  $\mu$ m and the laser power on the transparent single grains corresponded to 8 mW. Selected spectra were recorded with lower laser power of 1 and 4 mW on the sample to ensure that the samples were not modified by the measurements.

## RESULTS

### An<sub>50-57</sub> Plagioclase Shocked at Initial Temperatures of 77 K (–196 °C) and 293 K (20 °C)

Troctolite disks containing millimeter-size An<sub>50-57</sub> plagioclase grains were experimentally shocked at initial temperatures of 293 and 77 K and at pressures of 28 and 32 GPa, covering the approximate pressure range for maskelynitization (Ostertag 1983; Stöffler et al. 1986; Heymann and Hörz 1990) (Table 1). Shock experiments at higher pressures of 36 and 45 GPa were conducted only at 77 K. Transmitted light microscopy with crossed polarizers on thin sections showed that

plagioclase from the 28 GPa experiments was almost fully isotropic with only few birefringent domains observable in polished grain mounts. All plagioclase from the 32, 36, and 45 GPa shock experiments were completely isotropic.

The RI and chemical composition of the shocked plagioclase in the troctolite are plotted in Fig. 1. In accordance with microscopic observations, the data points for the 28 GPa experiments conducted at initial temperatures of 77 and 293 K were transitional between the fields for fully and partly isotropic material. The 28 GPa experiments were performed to obtain partly birefringent material, but only a few birefringent domains were observed in the recovered material. The material recovered from the 28 GPa shock experiments at initial temperatures of 77 and 293 K displays identical optical properties and clusters around the 30 GPa isobar of the calibration diagram (Fig. 1).

The data from the 32 GPa experiments at 77 and 293 K display a similar trend. In the calibration diagram (Fig. 1), data fall along isobars indicating slightly (2–3 GPa) higher shock pressures. Maskelynite with identical optical properties was recovered from the 32 GPa experiments conducted at initial temperatures of 77 and 293 K. The 36 GPa shock experiment at initial temperatures of 77 K produced maskelynite similar to that produced in the 32 GPa experiments at 77 and 293 K with data points plotting between the 34 and 40 GPa isobars in the calibration diagram (Fig. 1). The shock experiments at 45 GPa and initial temperatures of 77 K produced material with optical properties similar to those of a thermal mineral glass. The 45 GPa data in Fig. 1 show significantly less scatter than data from shock experiments between 28 and 36 GPa.

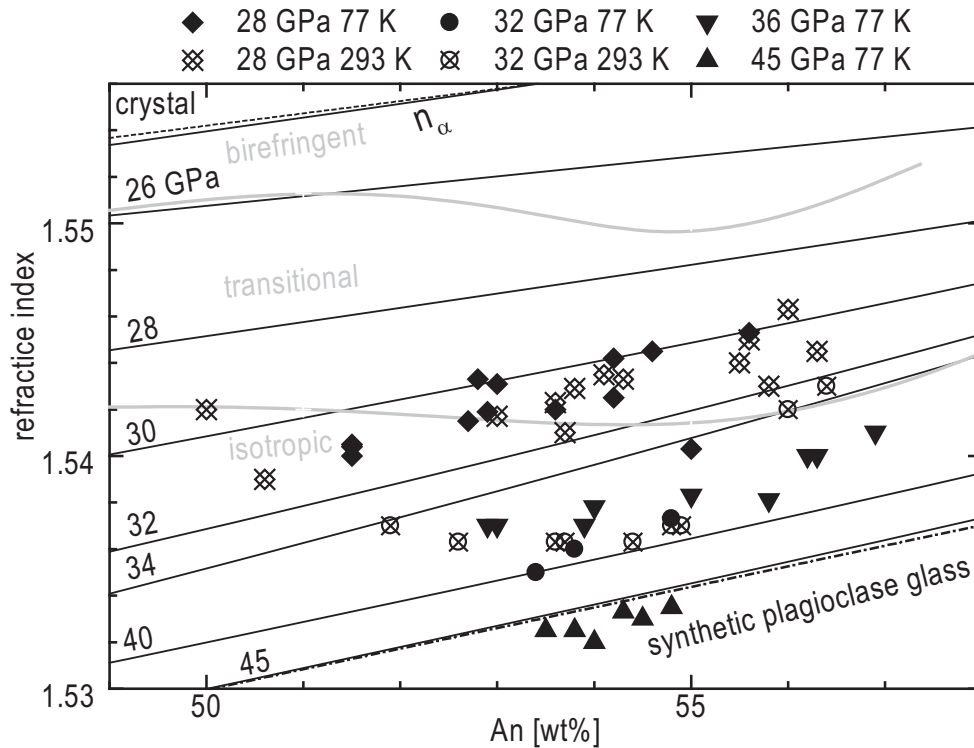


Fig. 1. Refractive index (RI) versus chemical composition (anorthite content = An) of plagioclase samples experimentally shocked to different pressures and at different initial temperatures of the experimental setup. Refractive indices of plagioclase crystals (for  $n_o$ ) and their synthetic (thermal) glasses from Tröger (1982). Solid black lines represent isobars for shock pressures that produce from plagioclase materials with the according optical properties; i.e., showing the indicated RI and being either birefringent, partly birefringent, or fully isotropic. The gray lines separate the fields of birefringent and partly birefringent plagioclase from fully isotropic maskelynite (redrawn after Stöffler et al. 1986).

#### **An<sub>94</sub> Plagioclase Shocked at an Initial Temperature of 293 K (20 °C)**

The influence of the chemical composition on the formation of maskelynite is investigated by room temperature experiments using the Grass Valley gabbro that contains next to pyroxene a homogeneous, Ca-rich plagioclase of An<sub>94</sub>. Visual inspection of the unshocked, polished rock disks reveals millimeter-size grains of dark green pyroxene and up to centimeter-size white plagioclase, with both minerals displaying an intense luster. These whole rock disks were shocked to pressures between 20 and 41.5 GPa and at initial temperatures of 293 K (20 °C). In the 20 GPa-shocked sample, both pyroxene and plagioclase are strongly fractured and the minerals appear milky, pale, and without luster. In contrast, plagioclase recovered from the 24, 28, 36 and 41.5 GPa shock experiments appears more coherent, significantly less fractured, and of increasing transparency. Pyroxene appears highly fractured and without luster during visual inspection.

The RI of unshocked and shocked plagioclase An<sub>94</sub> is shown in Fig. 2. The unshocked and the 20 GPa shocked plagioclase grains are birefringent and both

have an RI of 1.586 that is slightly higher than the RI of 1.584 for crystalline An<sub>94</sub> plagioclase as reported by Tröger (1982). At shock pressures of 24 GPa, plagioclase was almost fully transformed to isotropic maskelynite. In maskelynite formed at pressures of 24, 32, and 36 GPa, the RIs are 1.584, 1.582, and 1.576, respectively. These RI values are within the range of an isochemical crystalline plagioclase (Fig. 2), and within the range of RI determined from lunar maskelynite (An<sub>84-96</sub>) in Apollo 11 soil samples (von Engelhardt et al. 1970). The 41.5 GPa shock experiment produced amorphous material with an RI between 1.586 and 1.590, which is higher than the RI of the unshocked starting material (Fig. 2). For comparison, Fig. 2 displays literature data from Reynard et al. (1999) for synthetic anorthite glass (An<sub>100</sub>) shocked to pressures of 17–41 GPa together with the RI of An<sub>94</sub> maskelynite from the 1170 K heating experiment (see next paragraph).

#### **Heating Experiments with Maskelynite of Different Chemical Compositions**

Maskelynite produced in the 24 GPa shock experiment from Ca-rich plagioclase (An<sub>94</sub>) and in a

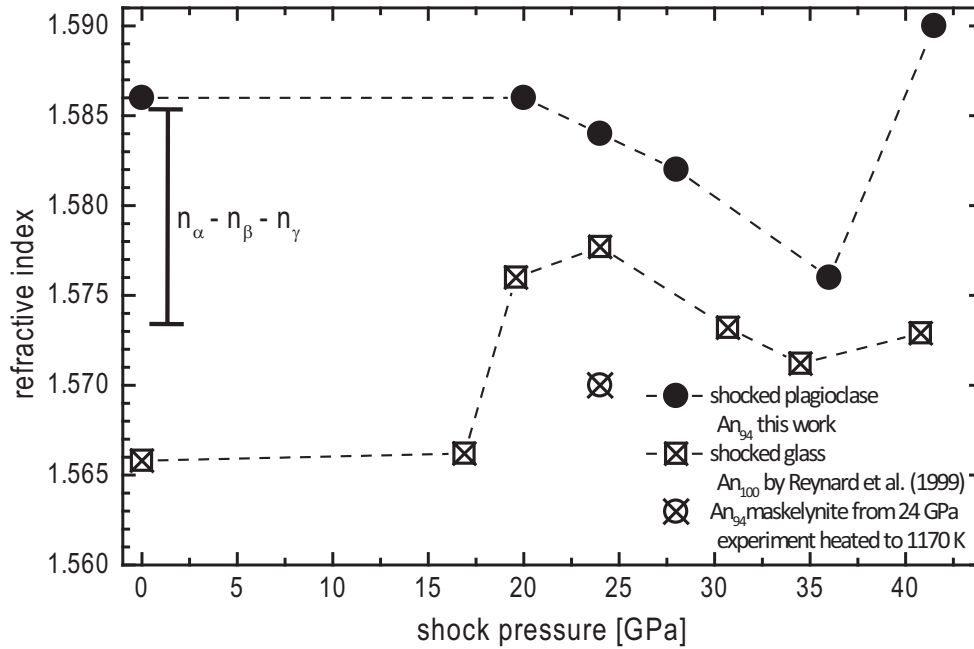


Fig. 2. Refractive index (RI) of experimentally shocked plagioclase (An<sub>94</sub>—this work) and synthetic glass (An<sub>100</sub>—Reynard et al. 1999). Additionally displayed is the RI of a plagioclase shocked to 24 GPa and then heated in an oven at temperatures of 1170 K. It has an RI identical to those of an isochemical glass (Tröger 1982). The range of refractive indices ( $n_\alpha$ ,  $n_\beta$ ,  $n_\gamma$ ) for crystalline plagioclase with An<sub>94</sub> as given by Tröger (1982) is indicated on the left.

32 GPa shock experiment from plagioclase of intermediate composition (An<sub>50-57</sub>) were used in 50 K step heating experiments at temperatures from 570 to 1170 K (Fig. 3). Up to temperatures of 770 K the RI of the An<sub>94</sub> sample remains unaffected. The RI slightly decreased after being exposed to 820 K for 5 min, and it remained unchanged until 870 K; it was then followed by a small decrease after being heated to temperatures of 920 and 1020 K. Further increase in the temperatures to 1170 K finally produced a sample with an RI of 1.570, equal to the RI of an isochemical thermal glass. For the relatively Na-rich An<sub>50-57</sub> sample shocked to 32 GPa, the initial RI of 1.538 was unaffected up to 920 K. At still higher temperatures, the RI of the An<sub>50-57</sub> sample decreased in a similar fashion to that of the An<sub>94</sub> sample (Fig. 3). At temperatures of 1170 K, the An<sub>50-57</sub> samples acquired an RI of 1.532, the value of an isochemical thermal glass.

#### Raman Spectroscopy on Plagioclase of Different Chemical Composition Shocked to Different Pressures

Representative Raman spectra from plagioclase and maskelynite with An<sub>55-57</sub> are displayed in Fig. 4. The unshocked sample shows spectra characteristic of plagioclase, including intense bands at 480 and 511 cm<sup>-1</sup>, and a relatively smaller band at 566 cm<sup>-1</sup> (Mernagh 1991; Freeman et al. 2008). With increasing shock pressure the 566 cm<sup>-1</sup> band increases relative to the 480

and 511 cm<sup>-1</sup> bands, and these three bands form a broad hump. The 797 cm<sup>-1</sup> band decreases and broadens and all bands below 410 cm<sup>-1</sup> disappear. In addition, a prominent broad hump between 900 and 1150 cm<sup>-1</sup> develops. In a few locations of the 28 GPa samples shocked at temperatures of 77 and 293 K, the collected spectra acquire narrow bands at 486 and 511 cm<sup>-1</sup> on the broad hump, thus indicating a few regions with incomplete amorphization. The typical Raman spectra of maskelynite are characterized by two broad humps between 450 to 600 and 950 to 1150 cm<sup>-1</sup>. With increasing shock pressure and decreasing RI, the shape of the broad hump between 450 and 600 cm<sup>-1</sup> changes as the relative contribution of the 570 cm<sup>-1</sup> band decreases (Fig. 4).

Raman spectra from plagioclase and maskelynite with An<sub>94</sub> reveal a similar evolution (Fig. 5). Unshocked plagioclase and plagioclase shocked to 20 GPa show nearly identical Raman spectra with distinct bands characteristic of plagioclase. The 24 GPa sample displays two broad humps between 450 to 600 and 950 to 1150 cm<sup>-1</sup> indicative of maskelynite. Raman spectra from the An<sub>94</sub> samples shocked to 41.5 GPa display 570/500 cm<sup>-1</sup> band intensity ratio (the relative contribution of the 570 cm<sup>-1</sup> band) similar to that from spectra of isochemical maskelynite produced by lower shock pressures (Fig. 5). Notably, An<sub>94</sub> maskelynite produced in the 41.5 GPa shock experiment has an RI substantially



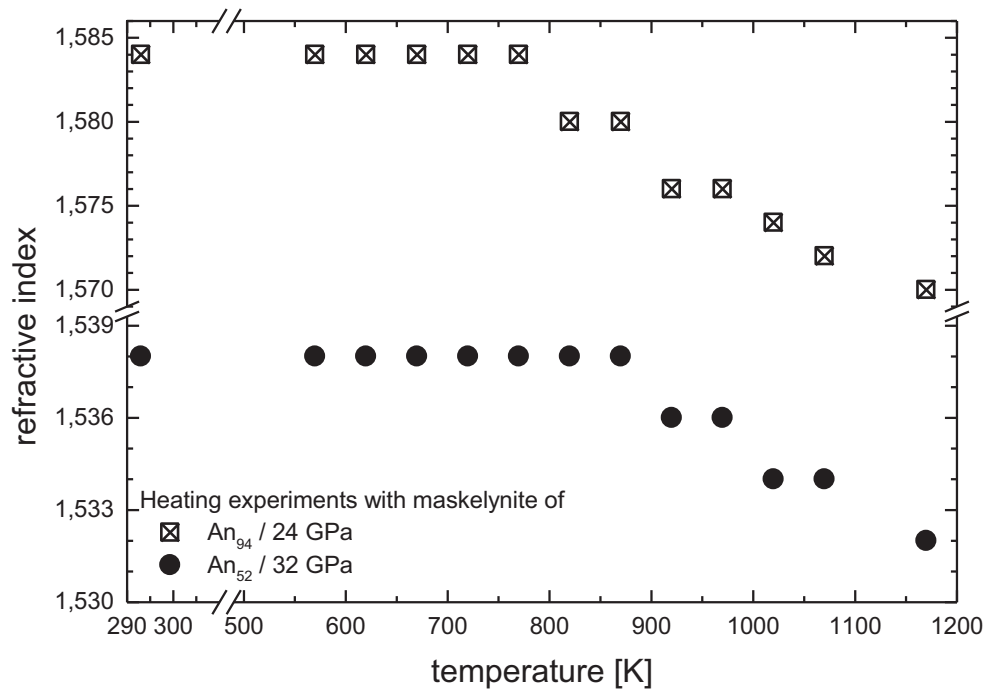


Fig. 3. Refractive index versus temperature determined with heating experiments using the shock experimentally produced maskelynite.

higher than that of maskelynite produced at lower shock pressures, and higher than that of the starting material. For comparison, the Raman spectra of maskelynite with different chemical composition and maskelynite thermally annealed at temperatures of 1173 K (900 °C) for 5 min are displayed in Fig. 6. The position of the 950–1150  $\text{cm}^{-1}$  hump depends on the chemical composition of maskelynite and the maximum shifts from 1100  $\text{cm}^{-1}$  in Na-rich maskelynite to  $\sim$ 980  $\text{cm}^{-1}$  in Ca-rich maskelynite. With increasing An content of the maskelynite, the 570/500  $\text{cm}^{-1}$  band intensity ratio in the Raman spectra increases. In addition, the 570/500  $\text{cm}^{-1}$  band intensity ratio decreases with decreasing RI; i.e., the 570/500  $\text{cm}^{-1}$  ratio decreases from spectra obtained from the 28 GPa shock-produced maskelynite [An<sub>50-57</sub> and An<sub>94</sub>] to those in spectra from maskelynite grains exposed to temperatures of 1170 K for 5 min, as well as in the An<sub>50-57</sub> maskelynite recovered from the 45 GPa experiments. The Raman spectra from maskelynite provide information on the chemical composition and the RI and, thus, density of the amorphous material.

## DISCUSSION

### Shock Formation of Maskelynite

The observations described above on the experimentally produced maskelynite provide new information

on the effect(s) of: (1) temperature during shock compression, (2) plagioclase composition with solidus and liquidus temperatures increasing from 1391 to 1826 K along the albite-anorthite solid solution series, and (3) postshock temperature on the annealing of maskelynite.

1. If plagioclase were to transform into maskelynite from a hot shock melt that rapidly quenched during decompression (Chen and El Goresy 2000; El Goresy et al. 2013), the following would be expected: Compared to room temperature (293 K) experiments, higher shock pressures will be required in much colder (77 K) experiments to eventually reach the melting temperature of plagioclase. However, materials produced at identical shock pressures but temperatures of 77 and 293 K have identical RI and Raman spectroscopic properties (Figs. 1 and 4). These results are equivalent to experimentally shocked quartz single crystals that resulted in identical shock deformation effects at any given pressure, at variable initial temperatures of 77 and 548 K (Fritz et al. 2011). Thus, the formation of diaplectic glasses from both quartz (Fritz et al. 2011) and plagioclase (this work) at low temperatures (initial temperatures of 77–293 K for plagioclase and 77–548 K for quartz) is controlled by shock pressure only and is independent of temperature variations. This together with the

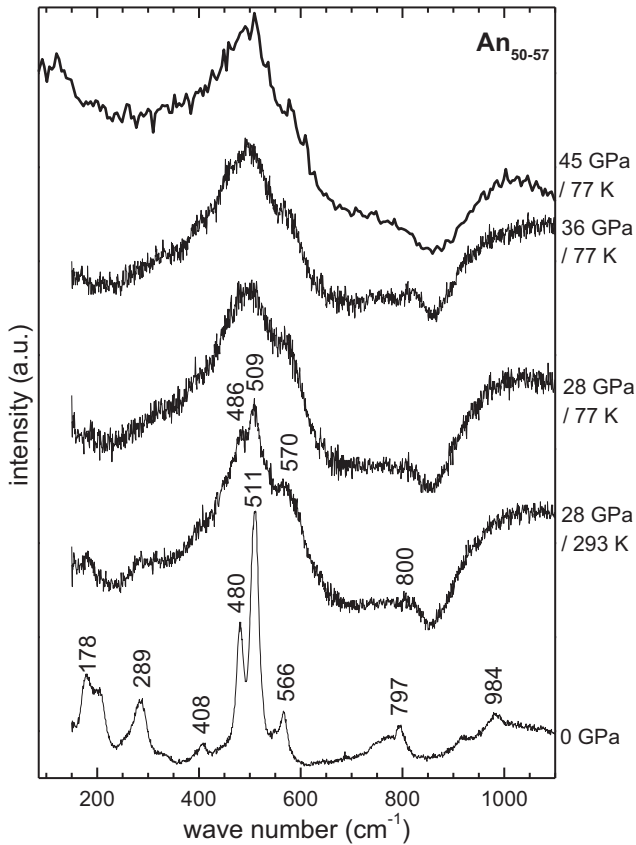


Fig. 4. Raman spectra of plagioclase with intermedium chemical composition ( $An_{50-57}$ ) shocked to different pressures and at different initial temperatures. Displayed are two types of spectra, each observed in samples from the 28 GPa experiments at 77 and at 293 K. Raman spectra from the 32 GPa samples (not shown) are very similar to the displayed spectra from the 36 GPa sample, and both samples have similar RI (see Fig. 1).

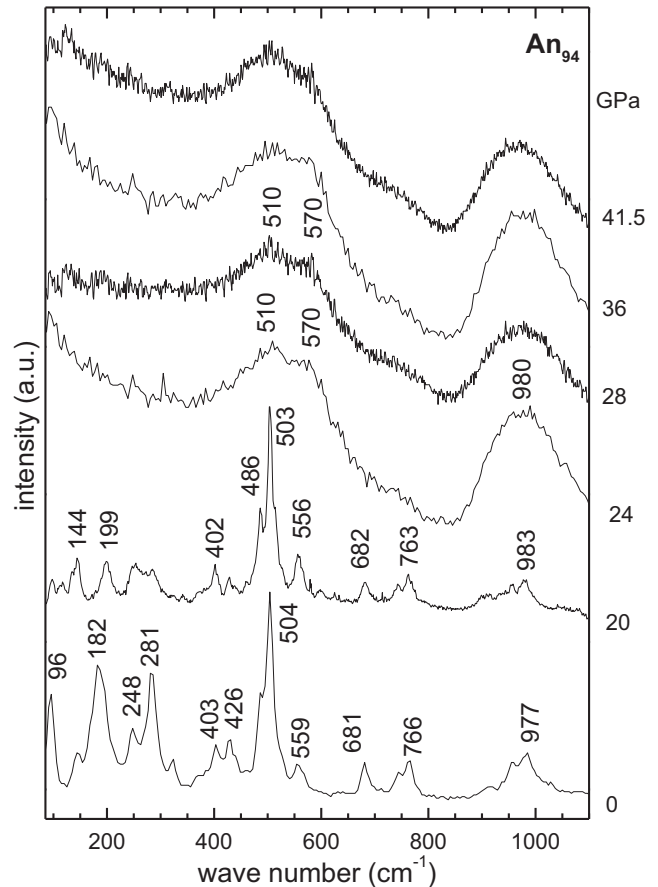


Fig. 5. Raman spectra of Ca-rich feldspar ( $An_{94}$ ) shocked to different pressures at initial temperatures of 293 K.

observation that plagioclase gradually transforms into maskelynite (Hörz and Quaide 1973) excludes diaplectic glass formation by thermal melting and advocates for a solid state amorphization. In static experiments, albitic plagioclase remains amorphous following decompression from pressures of 37 GPa at 440 K and from pressures of 31 GPa at 540 K (Tomioka et al. 2010). This indicates that plagioclase is more sensitive to elevated temperatures than quartz; i.e.,  $>440$  K for plagioclase versus  $>548$  K for quartz.

2. Compared to  $An_{50-57}$  plagioclase, the more Ca-rich  $An_{94}$  plagioclase has higher melting temperatures; i.e., at atmospheric pressure, the solidus of  $An_{50-57}$  and  $An_{94}$  is at 1500 and 1770 K, respectively. Consequently, it could be predicted that Ca-rich plagioclase should require higher shock pressures to melt during compression. However, consistent with

shock and static experiments (Ostertag 1983; Heymann and Hörz 1990; Daniel et al. 1997; Kubo et al. 2010; Johnson 2012), the opposite is observed, and the threshold pressure for the transformation decreases with increasing An-content (Fig. 7; Table 2). This trend can be explained by a pressure-induced failure of the plagioclase crystal lattice, and correlates with stronger Si-O bonds being replaced by weaker Al-O bonds toward An-rich compositions along the albite ( $Si/Al = 3/1$ ) to anorthite ( $Si/Al = 1/1$ ) solid solution series. Compared to Na-rich plagioclase in Martian meteorites or ordinary chondrites ( $An_{<75}$ ), the Ca-rich maskelynite in lunar and howardite, eucrite, and diogenite meteorites (e.g.,  $An_{>75}$ ) form at lower pressures and, thus can be associated with olivine and pyroxenes that display a lower degree of shock deformation effects.

3. The annealing experiments of maskelynite provide temperature limits for the postshock thermal history of a specimen at zero pressure. At atmospheric pressures and residing at temperatures of 770 and

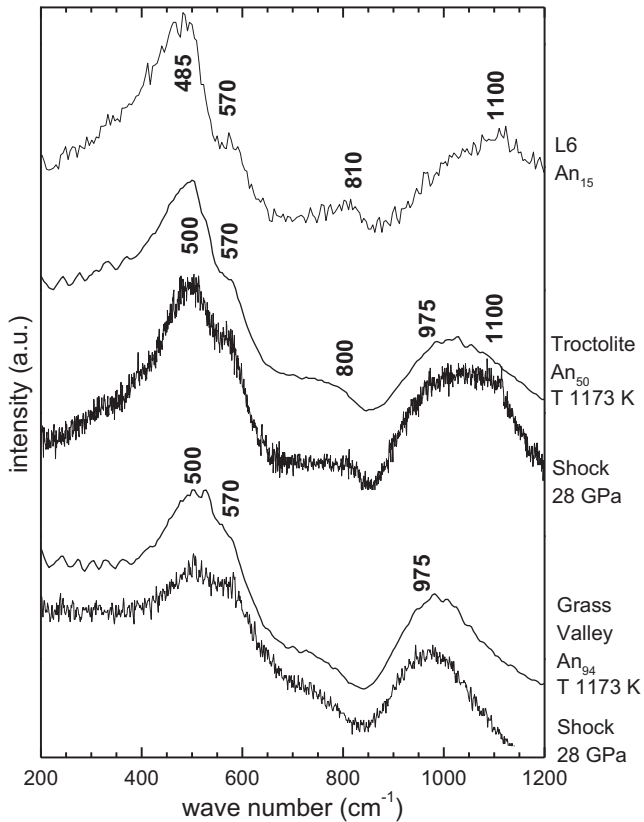


Fig. 6. Raman spectra of maskelynite with different chemical composition. The Grass Valley gabbro ( $An_{94}$ ) and a troctolite ( $An_{50-57}$ ) were experimentally shocked to 28 GPa. Raman spectra for these grains exposed for 5 min to temperatures of  $>1173$  K ( $900^{\circ}\text{C}$ ) are shown in addition. The L6 chondrite ( $An_{15}$ ) was naturally shocked. Note the variations in the 570/500  $\text{cm}^{-1}$  band intensity ratio and the position of the broad hump between 900 and 1150  $\text{cm}^{-1}$ .

900 K for 5 min at each temperature step, the RI of maskelynite starts to decrease (Fig. 2). The duration of our heating experiments is rather short compared with the cooling times in a natural environment, thus providing conservative upper limits for the thermal history of rocks containing maskelynite with the highest RI. In comparison, Duke (1968) used longer heating experiments on maskelynite from the Shergotty meteorite to show that it did not reach temperatures of 670 K following the shock event. Exposing maskelynite to temperatures of more than 1170 K for several hours to days eventually results in recrystallization to plagioclase (Ostertag and Stöffler 1982). The differences between the recrystallization of almost fully isotropic maskelynite and the crystallization of a thermal mineral glass are discussed by Arndt et al. (1982), indicating that maskelynite with high RI could retain a “memory” of the original crystal and, thus differs from a quenched

glass. The observed annealing behavior documents the gradual transition of plagioclase to maskelynite and thus supports a solid state transformation process. The effect of shock pressure and postshock temperature on the RI of maskelynite will be discussed in the following section.

The experimental data show that the shock formation of maskelynite is a function mainly of chemical composition and pressure and can occur at rather low temperatures during shock compression and decompression.

### Formation of Maskelynite with High and Low RI

This section discusses the effect of temperature and shock pressure on the RI of maskelynite by comparing the results of shock and thermal experiments (Figs. 1–3). Maskelynite with an RI close to the value of an isochemical thermal glass can be produced by increasing shock pressures to  $\sim 45$  GPa (Fig. 1) or by heating maskelynite with a relatively high RI produced from plagioclase shocked to lower pressures—e.g., to pressures in the range of 25–35 GPa (Figs. 1–3). Reynard et al. (1999) reported that the RI of the shock densified glass increases with a maximum density and an RI of 1.577 at 25 GPa (Fig. 2). At higher shock pressures, the recovered glass has a significantly lower RI value of 1.571 approaching the values of an isochemical thermal glass. Reynard et al. (1999) interpreted these experimental results as shock densification up to pressures of 25 GPa. These authors assumed that with increasing shock pressure, the higher postshock temperature causes thermal relaxation of the decompressed material. However, the experimental data presented here are somewhat inconsistent with the formation of low RI maskelynite only by postshock thermal relaxation of high RI maskelynite:

1. The 41.5 GPa shock experiment on  $An_{94}$  plagioclase resulted in maskelynite with an RI of 1.590 that is higher than those for unshocked plagioclase and maskelynite produced at shock pressures of 36 GPa (Fig. 2). Amorphous plagioclase ( $An_{96}$ ) with a very high RI of 1.62 was previously reported from static pressure experiments at temperatures of 293 K (Williams and Jeanloz 1989). Notably, Reynard et al. (1999) also recovered shocked synthetic glass ( $An_{100}$ ) with an RI that was slightly higher in the 41 GPa compared with the 35 GPa shock experiment (Fig. 2). These observations are inconsistent with a postshock temperature-induced reduction in the RI.
2. The 32 GPa shock experiments at 77 and 293 K produced  $An_{50-57}$  maskelynite with identical optical properties and RI values and these RI values are



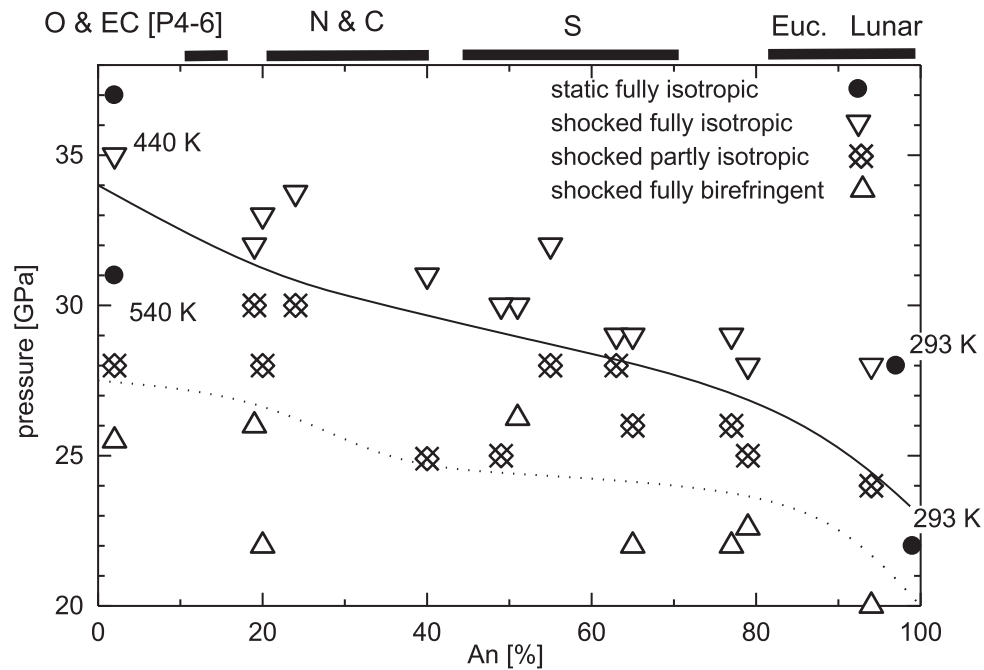


Fig. 7. Pressure versus anorthite content [An] for experimentally pressurized plagioclase. References are given in Table 2. Displayed are data from shock experiments conducted at initial temperatures of 293 K. The dotted and solid lines indicate the shock pressures above which crystalline plagioclase grains (birefringent) of a given chemical composition are transformed into grains composed of crystalline and amorphous regions (partly isotropic) and grains completely transformed into amorphous maskelynite (fully isotropic). The solid line indicating the minimum shock pressure to form  $An_{94}$  maskelynite approaches 24 GPa as this shock experiment produced almost fully isotropic material (this work). For static pressure experiments, the temperatures during maximum compression are shown in the graph, and all displayed data points represent isotropic maskelynite. Different pressure values for the armorphization of plagioclase with  $An_{96}$  and  $An_{100}$  recovered from static pressure experiments were reported by Williams and Jeanloz (1989) and Daniel et al. (1997), respectively. Not included in the graph is the partly isotropic  $An_{96}$  maskelynite recovered from 22 GPa experiments by Williams and Jeanloz (1989). The typical composition of plagioclase in different types of meteorites is indicated for lunar rocks (Moon), eucrites (Vesta?), the shergottites, chassignites, and nakhlites (Mars), as well as for equilibrated ordinary and enstatite chondrites (main belt asteroids). O & EC (P 4-6) = ordinary and enstatite chondrites of petrological type 4-6 (Kovach and Jones 2010); N & C = Nakhlites and Chassignites; S = Shergottites (Fritz et al. 2005a, 2005b); Euc = Eucrites (Mayne et al. 2009); Lunar (Wieczorek et al. 2006).

lower than those of  $An_{50-57}$  maskelynite from the 28 GPa experiments at 77 and 293 K. This suggests that the lower RI of the 32 GPa maskelynite compared with the 28 GPa maskelynite is related to pressure and not to temperature.

The formation of maskelynite requires shock pressures above the amorphization pressure of plagioclase with given composition. Maskelynite with high and low RI can be produced by shock pressure only. The RI of shock-produced maskelynite is sensitive to postshock thermal disturbances and thus provides an upper limit for the temperature of the rock after decompression (Fig. 3) (Duke 1968; Fritz et al. 2005b). In a naturally shocked rock, the shock pressures deduced from the RI of maskelynite can be compared to the degree of shock metamorphic overprint in olivine or pyroxene; i.e., compared to shock barometer that are less sensitive to postshock thermal disturbances.

### Comparison of Shock and Static Pressure Experiments

To date, most research groups working on shock metamorphism presume substantial differences between static and shock experiments (e.g., Stöffler and Langenhorst 1994; Sharp and De Carli 2006; Gillet et al. 2007). Shock and static experiments differ in time scales of compression and decompression, and may also differ in temperature during compression, including hydrostatic (no friction or internal stress) and nonhydrostatic (friction or oriented stress) conditions. Tomioka et al. (2010) and Kubo et al. (2010) concluded that amorphization of plagioclase occurs at significantly lower pressures (>10 GPa) in static compared to shock experiments. However, the present observations and literature data suggest that in both experimental settings the amorphization of plagioclase occurs at similar pressure and temperature conditions (Fig. 7; Table 2).

Table 2. Compilation of the shock and static pressure experiments on plagioclase with different chemical composition presented in Fig. 8.

Chemical composition	Temperature (K)	Type of experiment	References
An <sub>02</sub>	293	Shock	Johnson and Hörz (2003)
An <sub>19</sub>			Heymann and Hörz (1990)
An <sub>20</sub>			Ostertag (1983)
An <sub>24</sub>			Stöffler et al. (1986)
An <sub>40</sub>			Johnson (2012)
An <sub>49</sub>			Heymann and Hörz (1990)
An <sub>51</sub>			Stöffler et al. (1986)
An <sub>50-57</sub>			This work
An <sub>63</sub>			Gibbons and Ahrens (1977)
An <sub>77</sub>			Ostertag (1983)
An <sub>79</sub>			Johnson and Hörz (2003)
An <sub>94</sub>			This work
An <sub>01</sub>	440	Static	Tomioka et al. (2010)
An <sub>01</sub>	540		Tomioka et al. (2010)
An <sub>96</sub>	293		Williams and Jeanloz (1989)
An <sub>100</sub>	293		Daniel et al. (1997)

Data for An<sub>52</sub> plagioclase from in situ static pressure experiments (Kubo et al. 2010), and from shock experiments by Velde et al. (1989) and Kitamura et al. (1977) are not included here (see text).

Amorphization of Na-rich plagioclase (An<sub>01</sub>) was observed in static pressure experiments conducted at pressures of 37 GPa and temperatures of 293 and 440 K and at pressures of 31 GPa and temperatures of 540 K (Fig. 7) (Tomioka et al. 2010). Tomioka et al. (2010) compared these static threshold pressures of 31–37 GPa with the report of partly crystalline An<sub>01</sub> plagioclase recovered from 50 GPa shock experiments (Velde et al. 1989; Jaret et al. 2018). However, the results of Velde et al. (1989) and Jaret et al. (2018) are inconsistent with various shock recovery reports that observed complete amorphization of Na-rich plagioclase at much lower pressures (Fig. 7). For An<sub>01</sub> and An<sub>19</sub>, completely amorphous maskelynite was recovered from shock experiments at 35 and 32 GPa, respectively (Heymann and Hörz 1990; Johnson and Hörz 2003). These shock pressures of 35 to 32 GPa are identical to static pressures of 37 to 31 GPa at which amorphous

Na-rich plagioclase was recovered from static experiments by Tomioka et al. (2010).

For Ca-rich plagioclase (An<sub>100</sub>), amorphization pressures of 22 GPa (Daniel et al. 1997) were reported from static experiments at room temperature using a sample from a low metamorphic rock from Val di Fassa, Monzoni, Italy (Van Tendeloo et al. 1989). These pressure and temperature values are similar to the amorphization of plagioclase (An<sub>94</sub>) from the plutonic Grass Valley gabbro observed in 24 GPa shock experiments conducted at room temperature in this work. For volcanic anorthite (An<sub>96</sub>) from Miyake-zima, Izu Islands, Japan, the 22 GPa static experiments produced partly isotropic plagioclase and higher pressures of 28 GPa were required to produce fully isotropic maskelynite (Williams and Jeanloz 1989). Volcanic anorthite (An<sub>96</sub>) of a scoria fall from the Fugoppe, Hokkaido, Japan, experimentally shocked to 30 GPa displayed isotropic material including submicrometer thick lamellae of crystalline plagioclase, but still higher shock pressures were required to produce fully isotropic maskelynite (Kitamura et al. 1977). The apparent up to 6 GPa differences in pressure required to produce fully isotropic maskelynite from plutonic and volcanic anorthite might reflect the influence of the degree of Al/Si ordering. However, similar amorphization pressures are observed by comparing static and shock experiments for either plutonic or volcanic anorthite.

For an An<sub>52</sub> plagioclase of intermediate chemical composition, the amorphization pressures were reported during in situ static experiments where the shocked sample is characterized inside the pressure cell (Kubo et al. 2010). It seems that Kubo et al. (2010) determined the amorphization pressure during compression, because the amorphization of An<sub>52</sub> plagioclase occurred at 19 GPa and at temperatures of 473 K. For comparison, Daniel et al. (1997) report on static pressure experiments in which plagioclase (An<sub>96</sub>) becomes amorphous at 16 GPa during compression. However, Daniel et al. (1997) noted that higher pressures of 22 GPa are required to produce material that remains amorphous after decompression. Clearly, a meaningful comparison of static and shock experiments requires that both samples are studied after decompression to ambient conditions. The in situ data from Kubo et al. (2010) are therefore not included in Fig. 7.

Considering that the time duration of static (h) and microsecond shock recovery experiments vary by nine orders of magnitude, the differences in amorphization pressures for plagioclase of less than 3 GPa appear small. Hence, the results from microsecond shock recovery experiments can be directly applied to rocks exposed to shock pulse durations of milliseconds if not

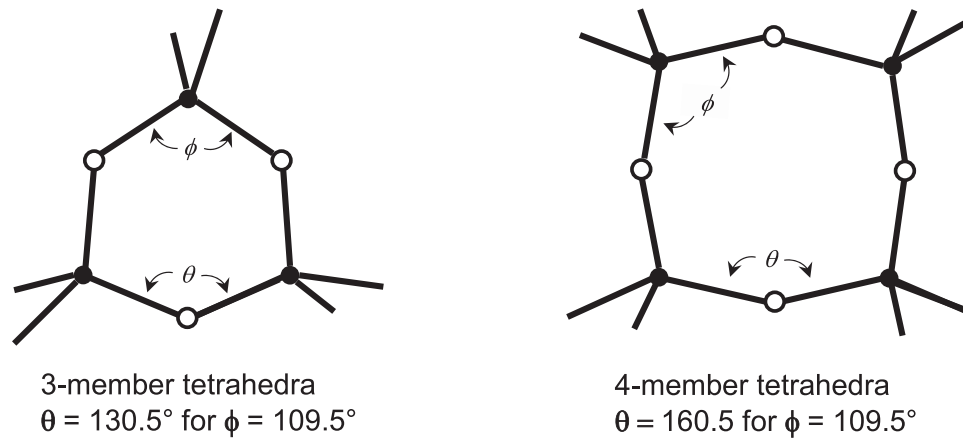


Fig. 8. Schematic drawing showing (a) three- and (b) four-membered planar rings of tetrahedra. Solid dots ( $\bullet$ ) represent T = Al or Si atoms and the open dots (O) represent O atoms. Values for the  $\theta$  angle were calculated by Galeener (1982) by assuming that the O-T-O angle ( $\Phi$ ) is tetrahedral ( $109.5^\circ$ ). Redrawn after Galeener (1982). In the maskelynite spectra (Figs. 4, 5, and 7), the Raman bands between  $400$  and  $600\text{ cm}^{-1}$  are attributed to the symmetric stretch of the oxygen atom along the line bisecting the T-O-T configuration ( $\theta$ ).

seconds during large-scale natural cratering events (Fritz et al. 2017).

Compared to shock and hydrostatic pressure experiments, the amorphization of plagioclase occurs at lower pressures in nonhydrostatic experiments (Daniel et al. 1997). A nonhydrostatic component, such as shearing along melt veins (Greshake et al. 2013), may explain the optical observations in several meteorites that isotropic maskelynite occurs locally near melt veins, yet the plagioclase remained birefringent elsewhere in these rocks (Binns 1967; Walton et al. 2016; Bischoff et al. 2019).

### Raman Spectral Properties

Raman spectroscopy documents substantial changes in the organization of the  $\text{TO}_4$  (T = Si, Al) tetrahedral units during the transformation of crystalline plagioclase to amorphous maskelynite (Heymann and Hörz 1990; Mernagh 1991; Reynard et al. 1999; Fritz et al. 2005b; Freeman et al. 2008). Raman spectra from unshocked plagioclase show strong bands in the  $450\text{--}600\text{ cm}^{-1}$  region that are related to the network of  $\text{TO}_4$  tetrahedra (Galeener 1982; Sharma et al. 1983; Matson et al. 1986; Mernagh 1991; Freeman et al. 2008). These tetrahedra form rings, and the number of ring members controls the bonding angle of bridging O atoms (Fig. 8). The motion of the oxygen atom along a line bisecting the T-O-T angle produces Raman bands in the spectral range of  $480\text{--}600\text{ cm}^{-1}$ . As the ring sizes decrease from four- to three-membered rings of tetrahedra, the Raman band positions increase from  $400$  to  $600\text{ cm}^{-1}$ . The variation of the  $570/510\text{ cm}^{-1}$  band intensity ratio serves as a measure for the relative

abundance of three- and four-membered rings of tetrahedra.

The increase in the  $570/510\text{ cm}^{-1}$  band intensity ratio associated with the transformation of plagioclase into maskelynite implies an increase in three-membered rings of tetrahedra at the expense of four-membered rings (Figs. 4, 5, and 8) (Reynard et al. 1999; Fritz et al. 2005b). Further increase in the shock pressure produces maskelynite with a lower RI index accompanied by a shift of the ring statistics toward larger ring sizes (Figs. 1 and 4). Identical Raman spectra obtained for maskelynite ( $\text{An}_{50-57}$ ) shocked to  $45\text{ GPa}$  and maskelynite shocked to  $28\text{ GPa}$  and then thermally annealed for  $5\text{ min}$  to  $1170\text{ K}$  (Fig. 6) show that these two approaches produce structurally similar material. The abundance of three-membered rings of tetrahedra (Fig. 6) increases with increasing An content, and this is likely related to the increasing Al/Si ratio.

The shock-induced loss of crystallinity is gradual as documented by several Raman spectral properties. Due to shock deformation of plagioclase, the  $480$  and  $510\text{ cm}^{-1}$  bands broaden, and narrow bands between  $650$  and  $800\text{ cm}^{-1}$  and below  $450\text{ cm}^{-1}$  disappear. The bands in the range of  $900\text{--}1170\text{ cm}^{-1}$  are related to T-O stretching vibrations of nonbridging O atoms attached to either Si or Al (Sharma et al. 1983; Matson et al. 1986). The absence of  $950\text{--}1150\text{ cm}^{-1}$  bands in unshocked plagioclase can be explained by the lack of nonbridging O atoms in a well-crystallized tectosilicate. Deformations and breaking of tetrahedral bonds during progressive amorphization lead to an increase in nonbridging O atoms attached to either Si or Al. Vibration of the nonbridging Si-O and Al-O units contributes to the broad hump at  $900\text{--}1150\text{ cm}^{-1}$ , and the shift of the maximum

position allows us to discriminate maskelynite of different chemical composition and associated lattice order (Fig. 6).

On a microscopic scale, this reorganization of the tetrahedral bonds is expressed by the transition from highly fractured, shocked plagioclase where micrometer-sized fragments would still retain the plagioclase crystal structure to a rather coherent mass of isotropic maskelynite. Therefore, the lack of fractures in maskelynite is not evidence for a hot, molten stage during maskelynitization (Chen and El Goresy 2000; El Goresy et al. 2013), but instead is better explained by substantial reorganizations in the tetrahedral network.

### Solid State Formation of Diaplectic Glasses and Seifertite in the Shergotty Meteorite at Pressures of 30 GPa

The thermal history during the transformation of plagioclase to maskelynite can be discussed using the pressure and temperature stability field of high pressure phases enclosed in maskelynite. In the Martian meteorite Shergotty (the maskelynite type specimen), maskelynite occasionally served as a shock pressure and temperature “container” for the formation of seifertite, a  $\alpha$ -PbO<sub>2</sub> type SiO<sub>2</sub> high pressure phase (Sharp et al. 1999; El Goresy et al. 2004). In a high temperature environment (>1000 K), seifertite is stable at pressures higher than the stability field of stishovite; i.e., depending on the temperature, it requires pressures in excess of 60–80 GPa (Sharp et al. 1999; Dubrovinsky et al. 2001; El Goresy et al. 2004). Lower temperatures inhibit the formation of stishovite, and seifertite can form from cristobalite at low temperatures (~480 K) and hydrostatic pressure of 25 GPa (Bläß 2013; Kubo et al. 2015; Cernok et al. 2017). Thus, the presence of seifertite in maskelynite of the Shergotty meteorite is consistent with a low temperature formation at shock pressures of 30 GPa as deduced from the shock effects in the rock-forming minerals, including maskelynite with a high RI (Stöffler et al. 1986; Fritz et al. 2005a).

## CONCLUSIONS

New shock experiments on plagioclase together with literature data show that the minimum pressure to form maskelynite decreases from 33 to  $\leq 24$  GPa along the albite to anorthite solid solution series, respectively. Shock recovery experiments at identical pressures and initial temperatures of 77 and 293 K produced maskelynite with identical optical properties in the present work. This is similar to the shock deformation of quartz, which resulted in identical optical properties at given shock pressure and initial temperatures of 77 and

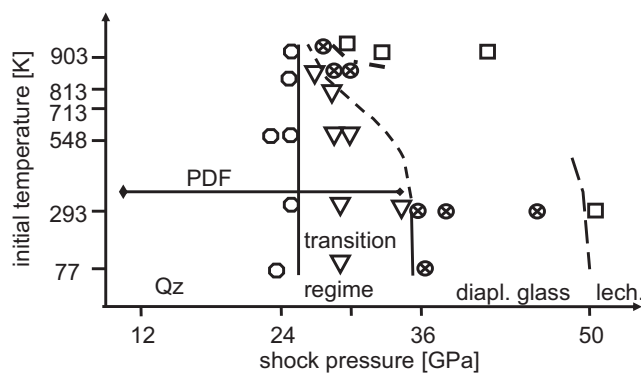


Fig. 9. Observed shock effects for target quartz of different initial temperature as a function of shock pressure. PDF = planar deformation feature; O = quartz; ∇ = quartz with reduced density, birefringence, and refractive index; ● = diaplectic glass; □ = lechatelierite. Redrawn after Fritz et al. (2011).

548 K (Fritz et al. 2011) (Fig. 9). Both results show that the formation of diaplectic glass is controlled by pressure, and it is rather independent of temperatures in the low temperature region—i.e., <540 K as inferred from the static pressure experiments by Tomioka et al. (2010). In addition, amorphization of plagioclase occurs at similar pressures ( $\pm 3$  GPa) in shock and static experiments as long as these experiments are conducted at the same temperatures, and the materials are characterized after decompression to atmospheric pressures. For quartz, fully amorphous material was recovered from static experiments at 293 K and pressures >30 GPa (Kingma et al. 1993) and from 34 GPa shock experiments at initial temperatures of 293 K (Stöffler and Langenhorst 1994). Microsecond shock recovery experiments on plagioclase and quartz produce the same materials as those recovered from static experiments at identical pressures and temperatures. This advocates that the properties of plagioclase, quartz, or cristobalite (see previous paragraph on seifertite) during shock compression can be studied using in situ observations of materials exposed to hydrostatic pressures at relevant temperatures.

It is established that the shock metamorphic effects of rock-forming minerals such as quartz, plagioclase, olivine, and pyroxene calibrated by shock recovery experiments can serve as a reliable and meaningful thermobarometer for the shock conditions in naturally shocked nonporous rocks (see also Fritz et al. 2017). It is notable that the minimum shock pressure for maskelynitization decreases with increasing An content, and thus is different for specific types of meteorites (Fig. 7).

For example in Martian meteorites, the shock pressure and temperature conditions (Fritz et al. 2005a,



2005b) were found to be highly consistent with a thermal history inferred from a variety of independent properties, including (1) the radiogenic  $^4\text{He}$  inventory (Schwenzer et al. 2008), (2) the magnetic properties of ALH84001 (Weiss et al. 2000); (3) the recrystallization of plagioclase in a decimeter-sized rock that cooled in space (Fritz et al. 2005a), (4) the formation of seifertite enclosed in maskelynite in the Shergotty meteorite (Sharp et al. 1999; Bläß 2013; Kubo et al. 2015; Cernok et al. 2017; this work), and (5) the formation and preservation of thermally unstable high pressure phases associated with shock-produced melt veins and pockets in some meteorites (Fritz et al. 2017; Tomioka and Miyahara 2017; Bischoff et al. 2019). The low temperature solid state formation of maskelynite does not reset the  $^{40}\text{Ar}$ - $^{39}\text{Ar}$  chronometer of rocks, as Ar diffusion requires a prolonged heating regime (Jessberger and Ostertag 1982; Fernandes et al. 2009; Shuster et al. 2010).

The formation of diaplectic glasses was compared with solid state or subsolidus transformations (Jaret et al. 2015) such as vitrification via an amorphous, highly coordinated structure during compression to a decompressed glass (Williams and Jeanloz 1989), via an increase in lattice defects similar to vitrification by radiation damage (Ashworth and Schneider 1985), or due to a gradual collapse of the lattice into domains that are eventually so small that they are beyond coherent X-ray diffraction of some 1–20 nm in size (Hörz and Quaide 1973). Due to their density and amorphous structure, the diaplectic glasses are often described as dense glasses. However, since they derive from a crystalline precursor due to pressure-induced collapse of the crystal lattice, these diaplectic glasses could better be viewed as inflated or shattered crystals, i.e., the term diaplectic (to destroy by striking) correctly describes the formation process.

*Acknowledgments*—Shock experiments were conducted during the research in the frame of the Helmholtz Alliance on Planetary Evolution and Life — WP3200. The sample disks and ARMCO container were prepared at the MfN by H.-R. Knöfler and H. Schneider. The gabbro was given to us by K. Ross and kindly provided by R. Wenk and B. Cassata (University of California at Berkeley). We thank Prof. Doug Crowe for providing rock sample B59-36-A from the Gilles O. Allard Economic Geology Collection, Athens, GA. Financial support for V. A. F. through the D. F. G. research grant FE 1523/3-1 and a Marie Skłodowska Curie Fellow, funded by the European Commission, HORIZON2020 Programme, grant number 749815 are acknowledged. We highly appreciate the detailed and constructive review by Fred Hörz and the Associate

Editor Gordon Osinski, who greatly helped to improve a previous version of this manuscript.

*Editorial Handling*—Dr. Gordon Osinski

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